

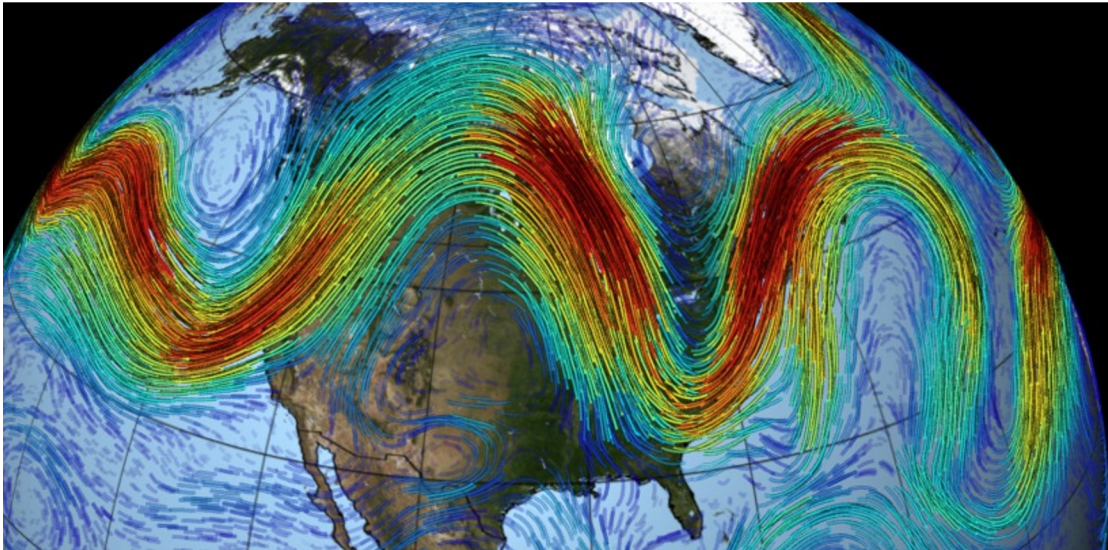
# NORTH ATLANTIC JET STREAM RESPONSE TO WARMING: TRENDS AND DRIVERS

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TRAMPAS meeting 24/10/2023

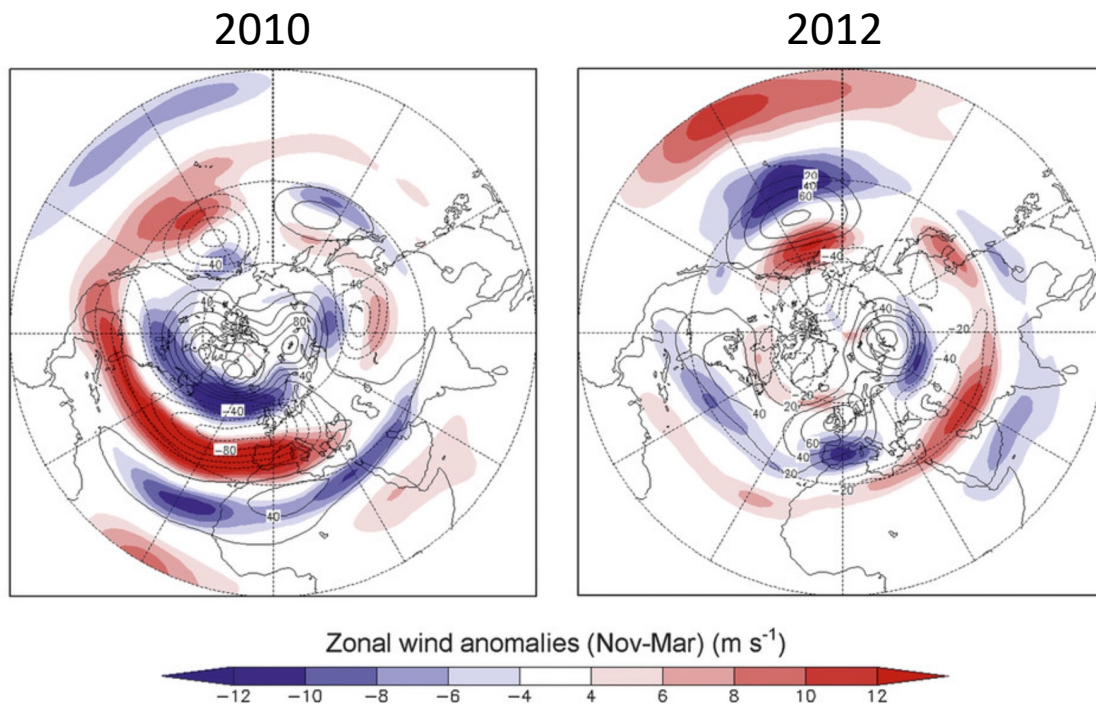
# INTRODUCTION

# Jet stream

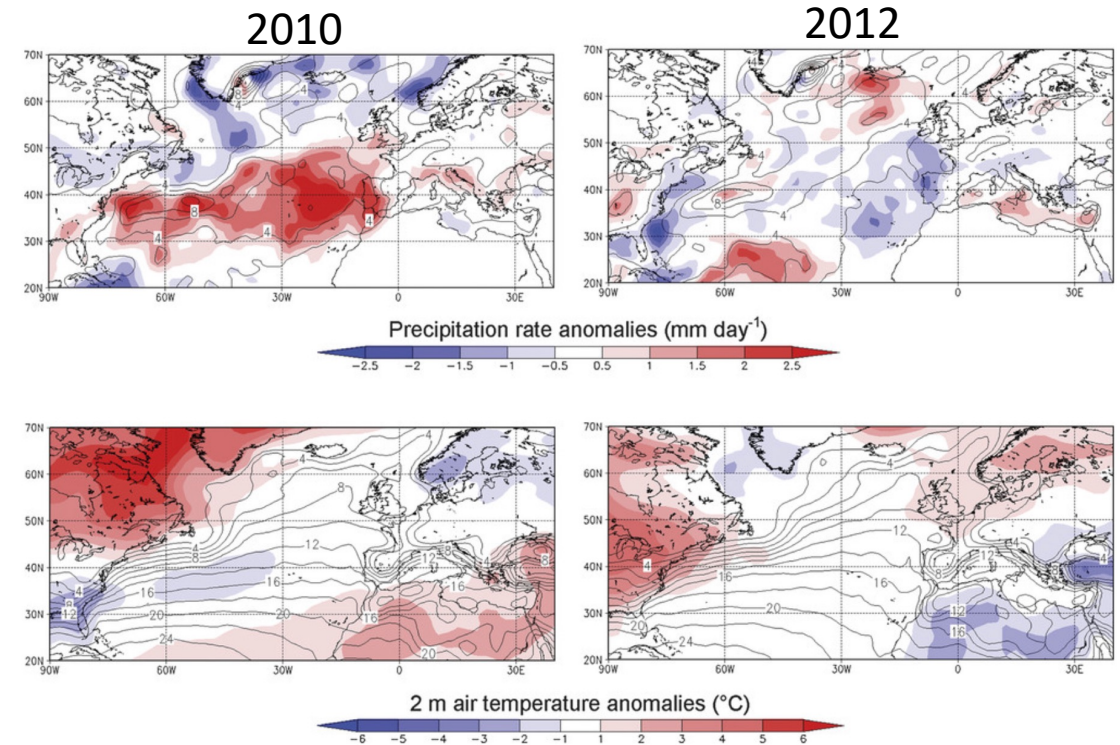


- Narrow and strong air currents flowing from west to east in both hemispheres
- Important influence on weather variability and extremes

# Exemple of impact of jet variability (Santos et al. 2013)



Anomalies of zonal wind at 250 hPa



Anomalies of precipitation (top) and T2 (bottom)

# Relevant processes

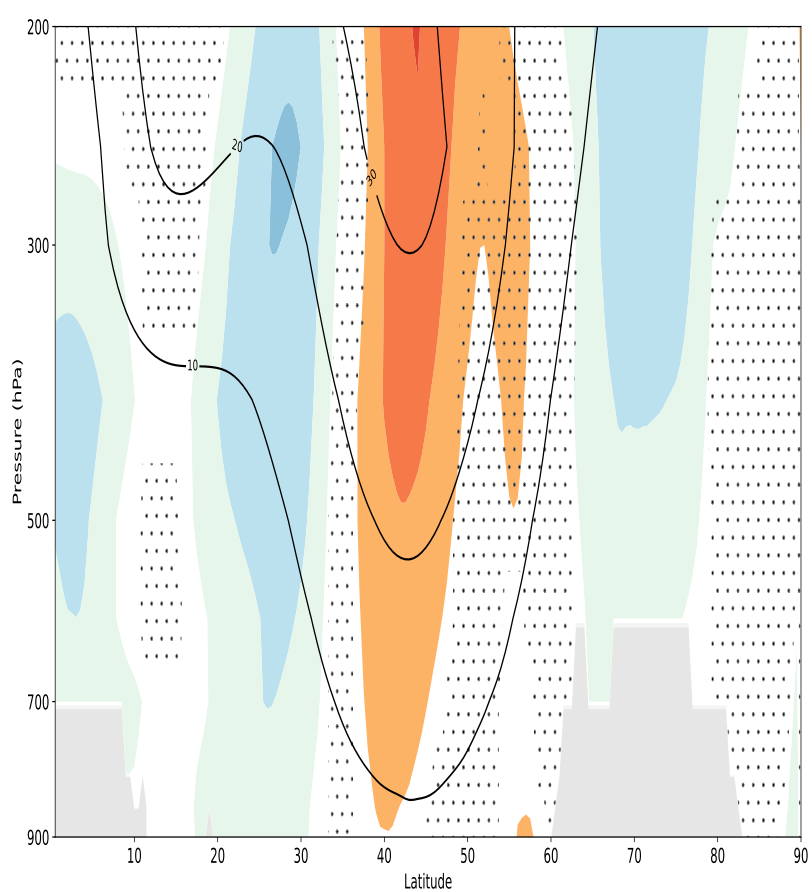
- Main mechanisms affecting the jet:
  - Areas of high baroclinicity -> development of Rossby waves -> eddy momentum flux convergence
- Local effect of diabatic heating
  - PV anomalies, which can intensify the jet

# Objective and data

- Analysis of jet stream trends (observed and projected) in winter (DJF) over the North Atlantic
- Understanding the mechanism leading to the observed trends
- Data
  - **ERA5** (1979-2022)
  - Idealised **2D frontal-geostrophic simulation**. Sawyer-Eliassen equation with diabatic heating
  - Idealised **aquaplanet simulations**. Control and uniform warming (4 K) + SST front
  - **CESM2 simulations**: 5 ensemble members, 1980-2100, SSP370 from 2015

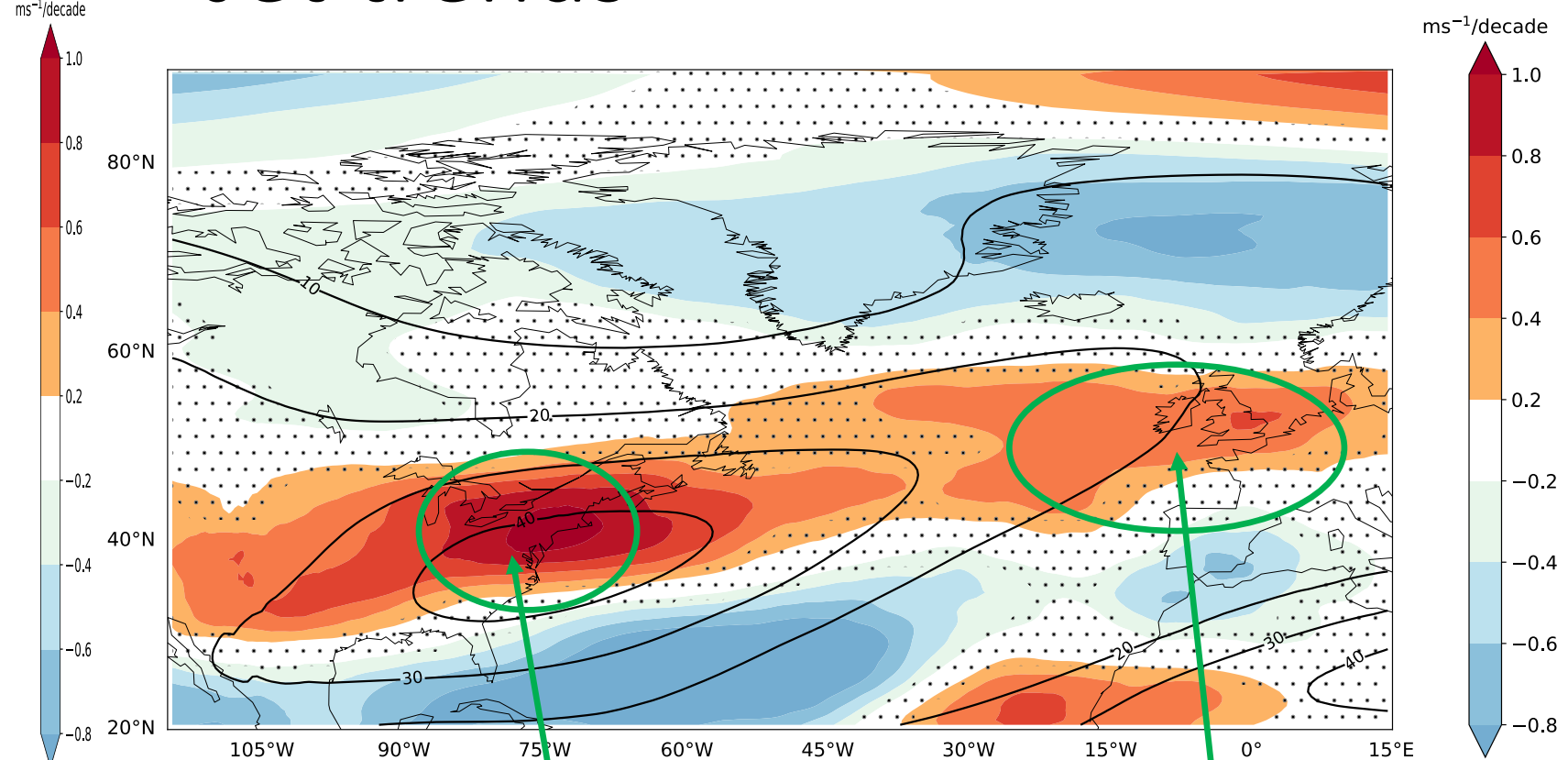
# Trends in ERA5

# Jet trends



**Zonal mean 80 – 15 W**

Shading: zonal wind speed trend  
 Black contours: climatological mean  
 Stippling: non-significant trends



**Trends at 250 hPa**

Intensification of the jet, slight poleward shift

Extension over Europe  
 Slight equatorward shift  
 downstream



# E vector

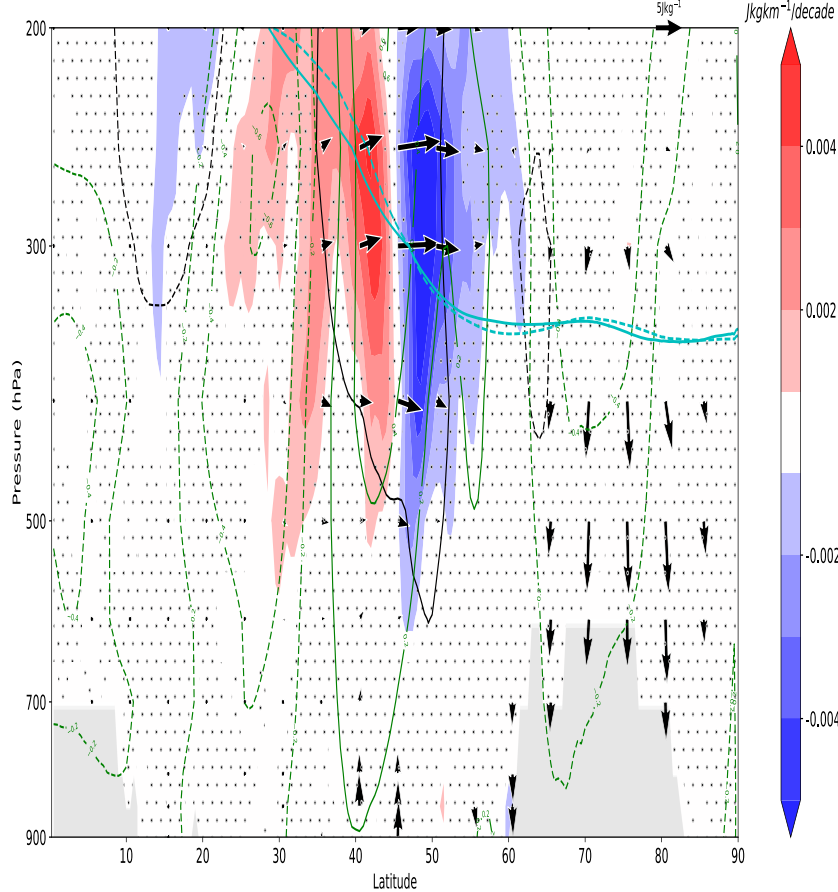
$$\mathbf{E} = \left( \overline{v'^2 - u'^2}, \overline{-u'v'}, \frac{f}{\partial\theta/\partial p} \overline{v'\theta'} \right)$$

Horizontal components: Eddy momentum flux

Vertical component: Eddy heat flux

E vector divergence -> Eddy momentum convergence

## E vector trends



**Zonal mean 80 – 15 W**

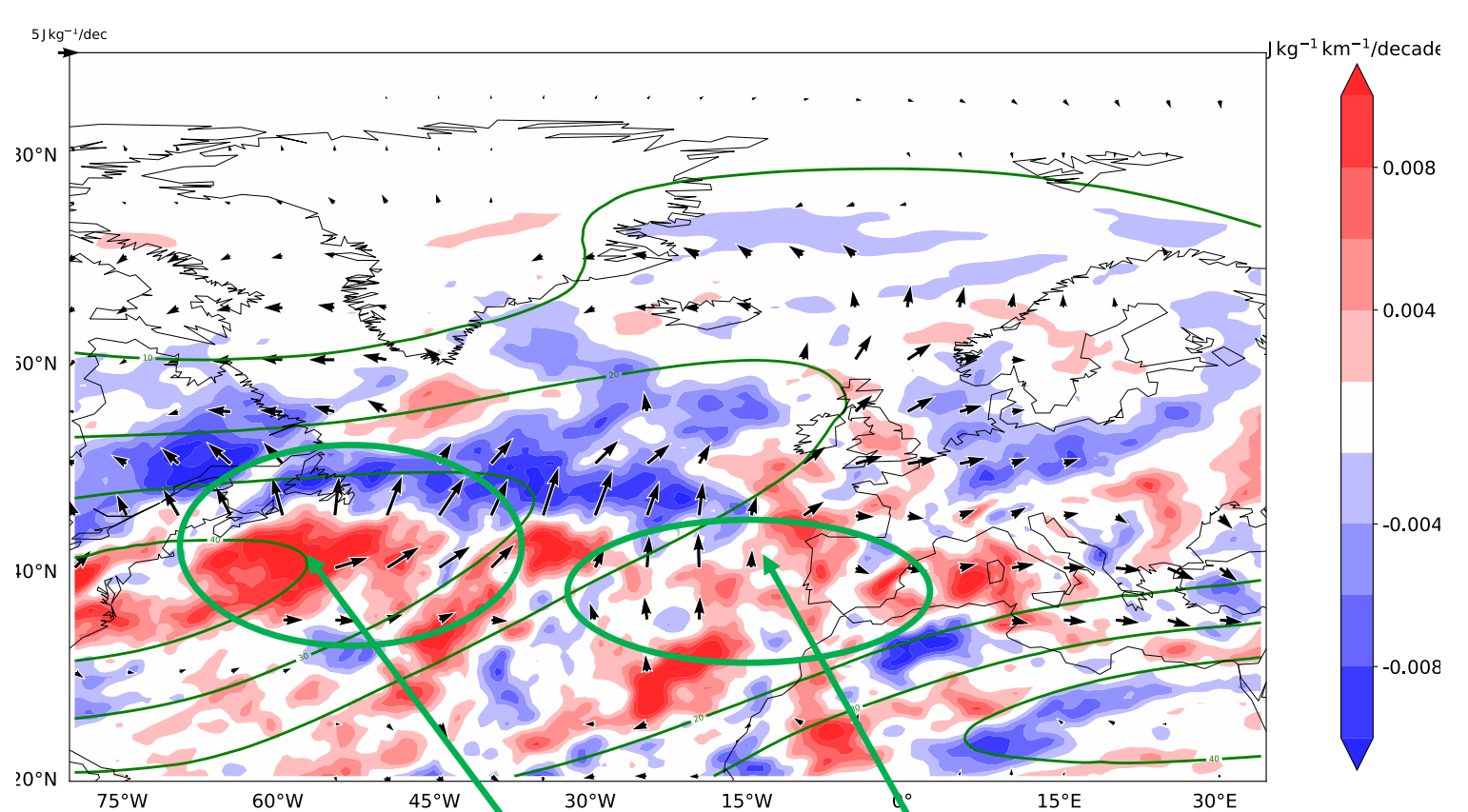
Shading: E vector divergence trend

Black contours: E vector divergence climatological mean

Green contours: Zonal wind trend (left)/ climatology (right)

Blue contours: Tropopause level (2PVU)

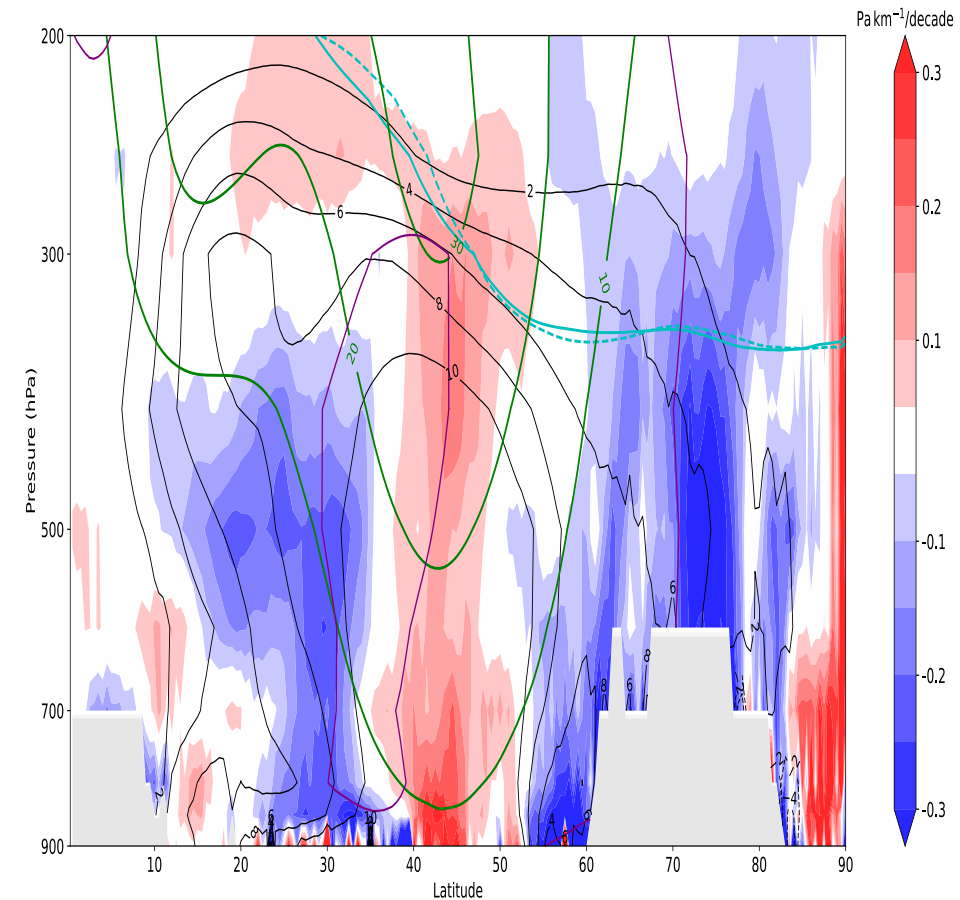
in first and last decade



**Trends at 250 hPa**

Increase of eddy momentum convergence  
around the jet core and equatorward downstream

# Slope trends



**Zonal mean 80 – 15 W**

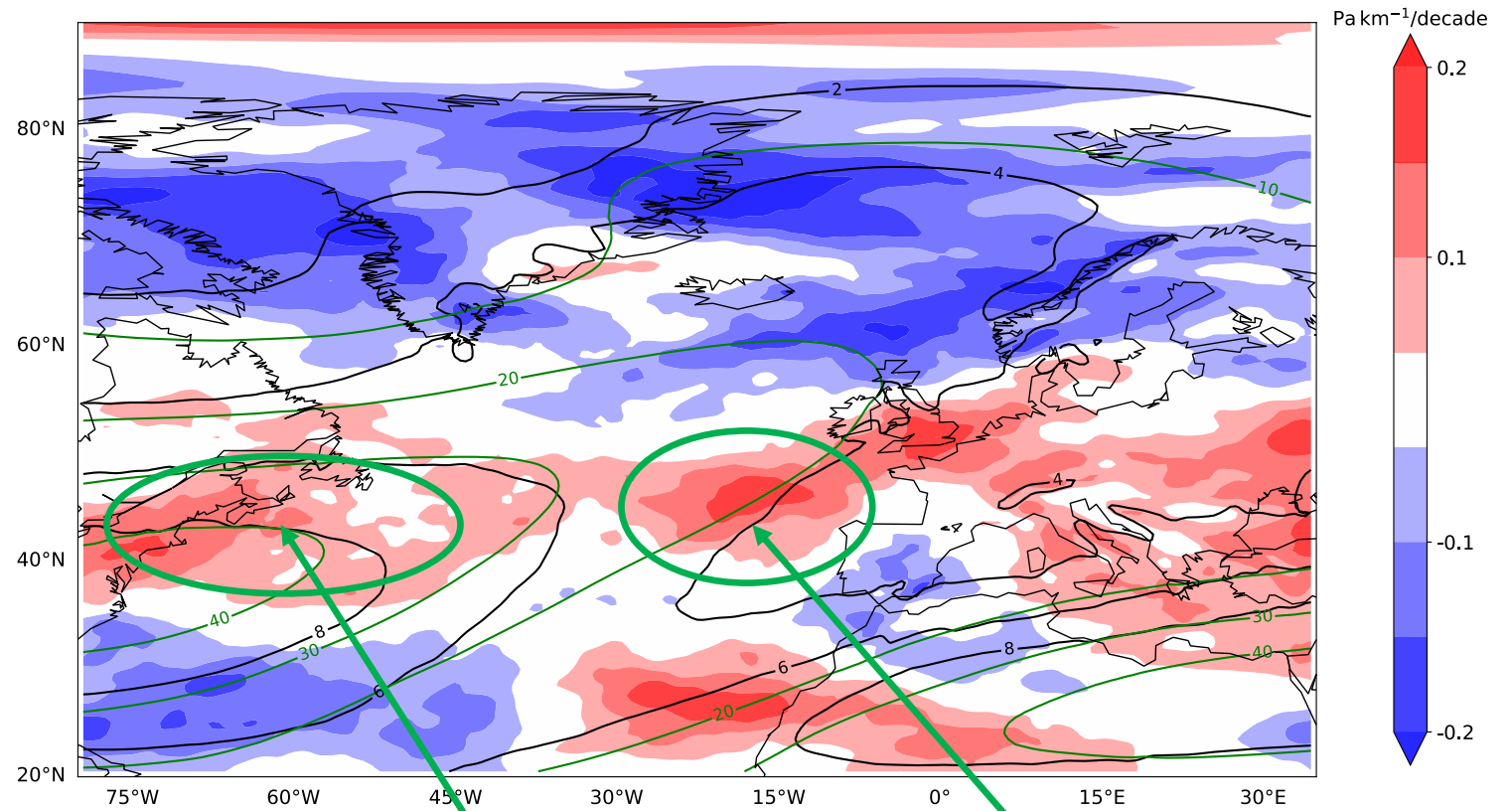
Shading: slope trend

Black contours: climatological mean

Green contours: Zonal wind speed climatology

Purple contour: Region of strongest potential temperature trend

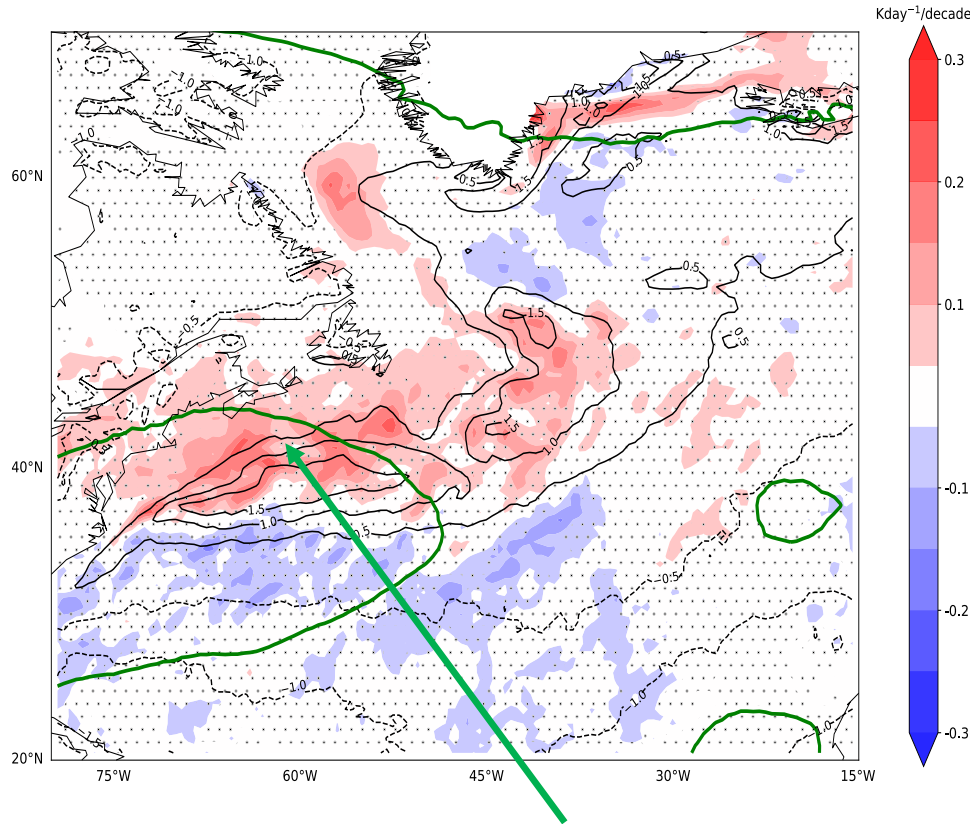
Blue contours: Tropopause level (2PVU) in first and last decade



**Average 200-500 hPa**

Slope increase around the climatological jet position  
over the Gulf stream and downstream

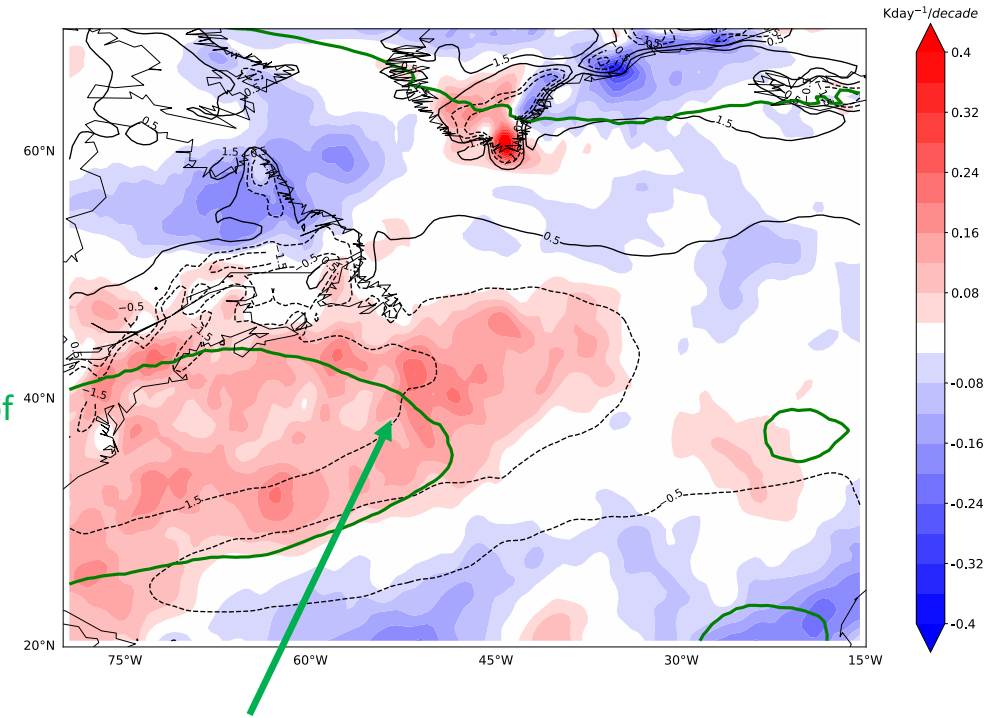
# Diabatic heating and advection trends



Increase over the Gulf Stream  
North of area of strongest potential temperature trend

Diabatic heating trend (300 – 850 hPa)

Shading: trends  
Black contours:  
climatological mean  
Green contour: Region of  
strongest potential  
temperature trend



Reduced cold advection over western NA ->  
reduced land-sea contrast  
Coincident with area of strongest potential temperature trend

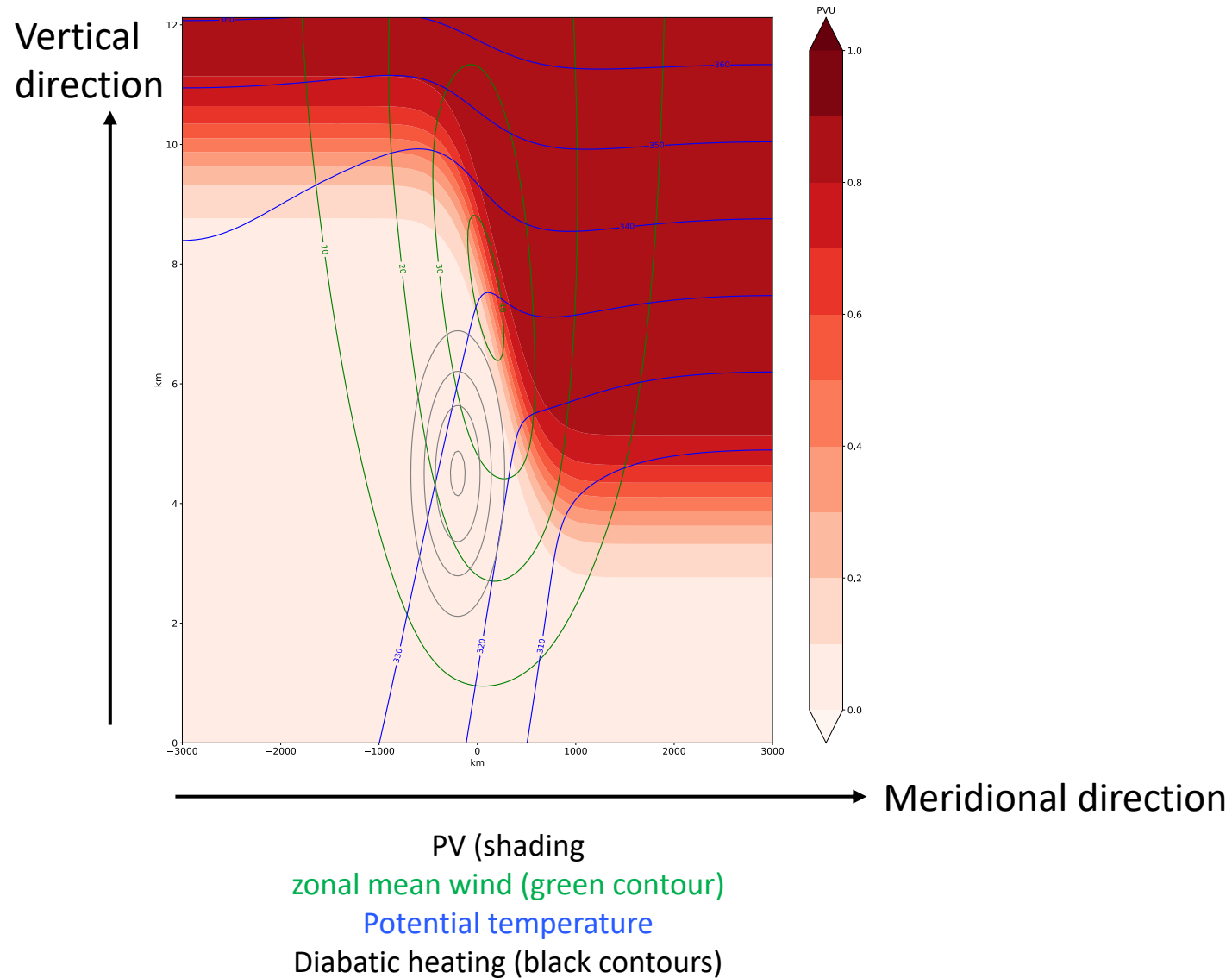
Advection trend (300 – 850 hPa)

Semi-geostrophic simulation:  
Local effect of diabatic heating

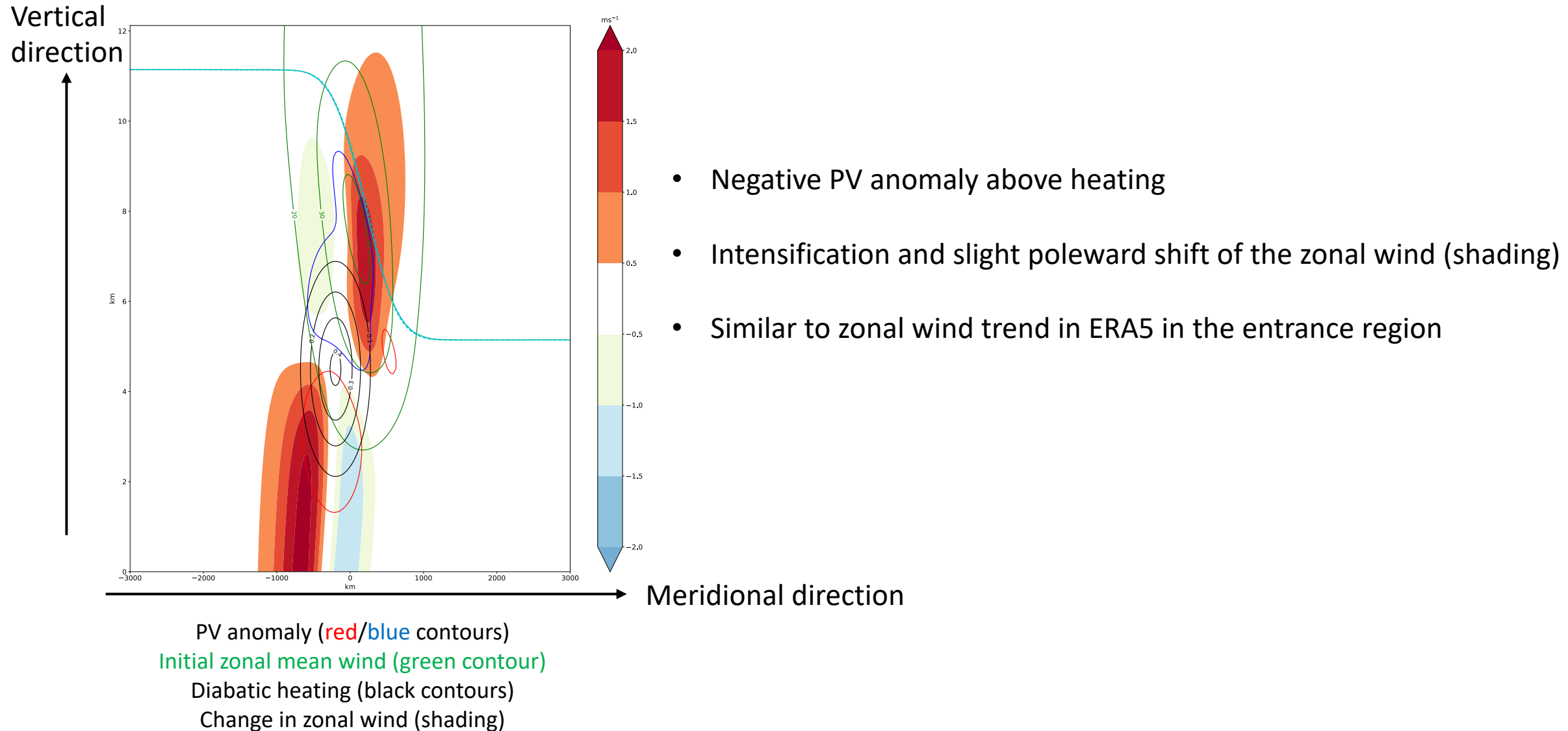
## 2D semi-geostrophic simulation

- Evolution of a zonally symmetric flow on an f-plane
- Zonal momentum  $m = u - fy$  and  $\theta$  advected on a y-z plane and modified by heating
- Meridional circulation obtained from thermal wind balance between  $m$  and  $\theta$
- The elliptic Sawyer-Eliassen equation for the streamfunction provided that PV is positive over the entire domain

# 2D semi-geostrophic simulation: Setup



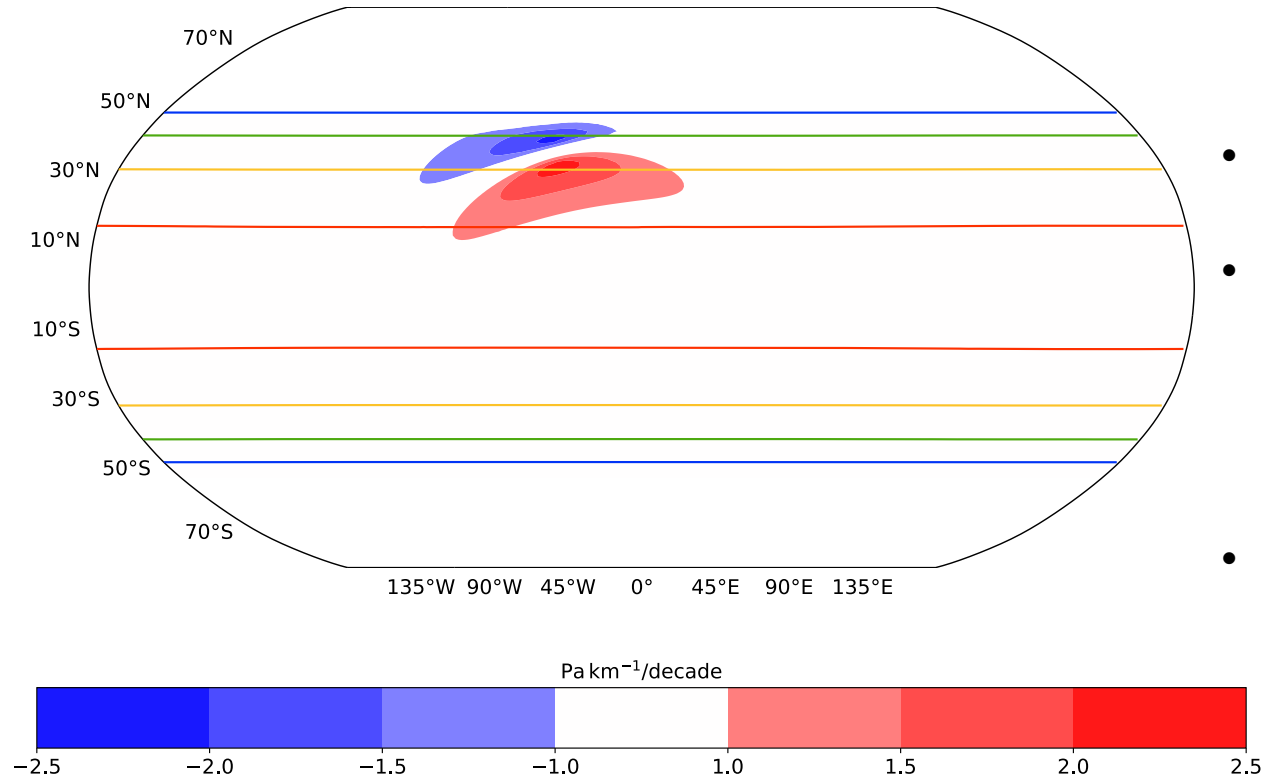
# 2D semi-geostrophic simulation: Results





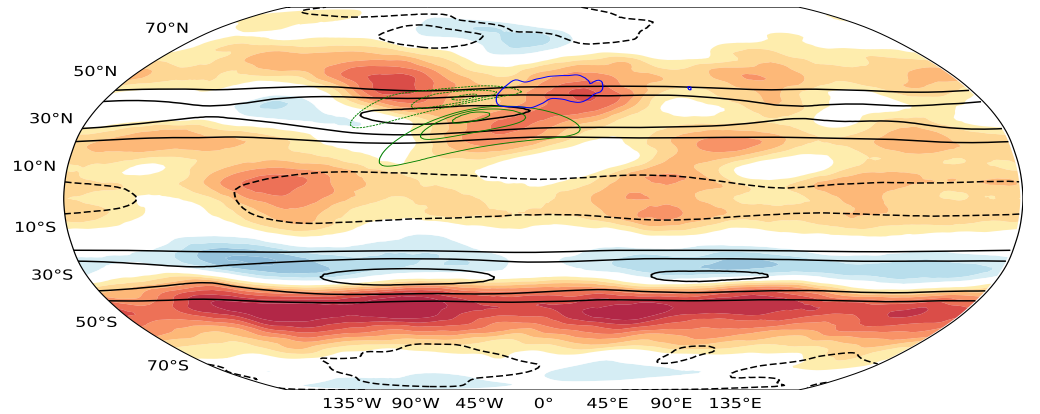
# Aquaplanet simulations: Eddy- mean flow feedback

# Aquaplanet simulations: Setup

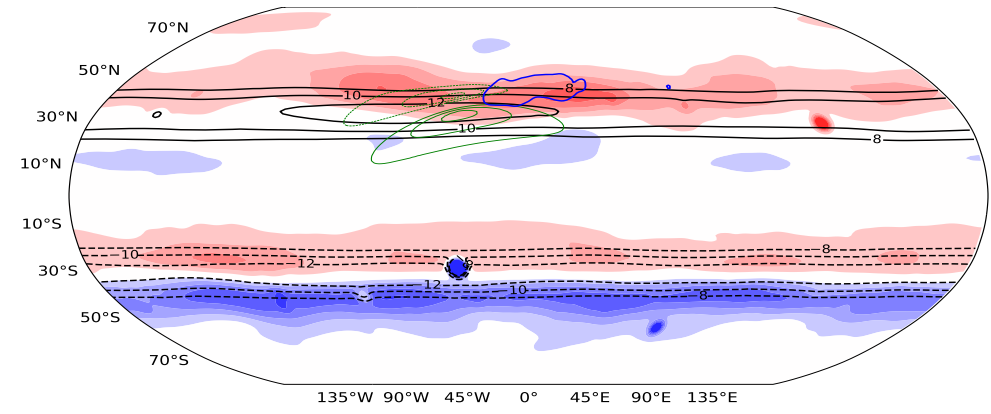


- ICON model v2.6.5, 80 km horizontal resolution
- **Control simulation:** Baseline SST profile + SST front (idealized land-sea contrast) at 30°W and different latitudinal positions (38, 39 and 43°N)
- **Warmed simulation:** Same setup with uniform warming (4K)

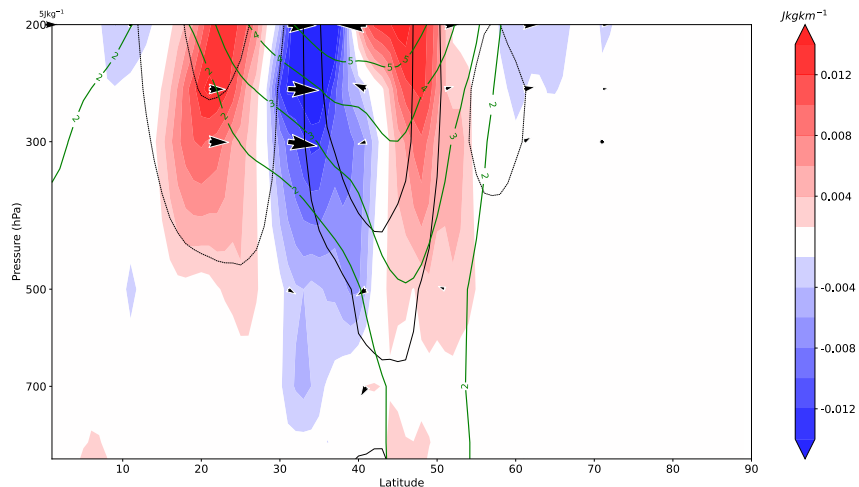
# Aquaplanet simulations (front at 39N)



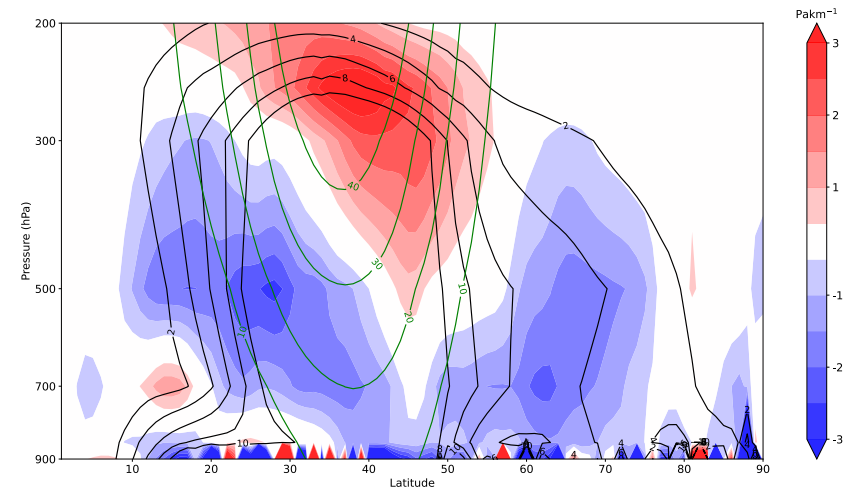
250-hPa zonal wind / diabatic heating (300-850 hPa)



Slope (200-500 hPa) / diabatic heating (300-850 hPa)

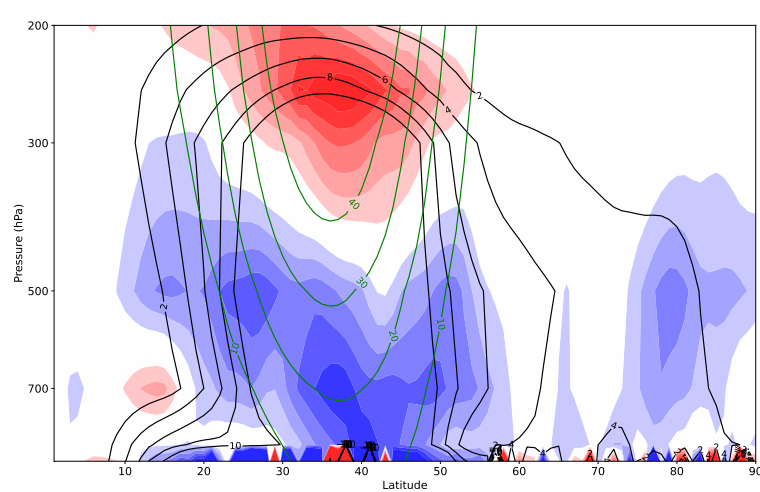
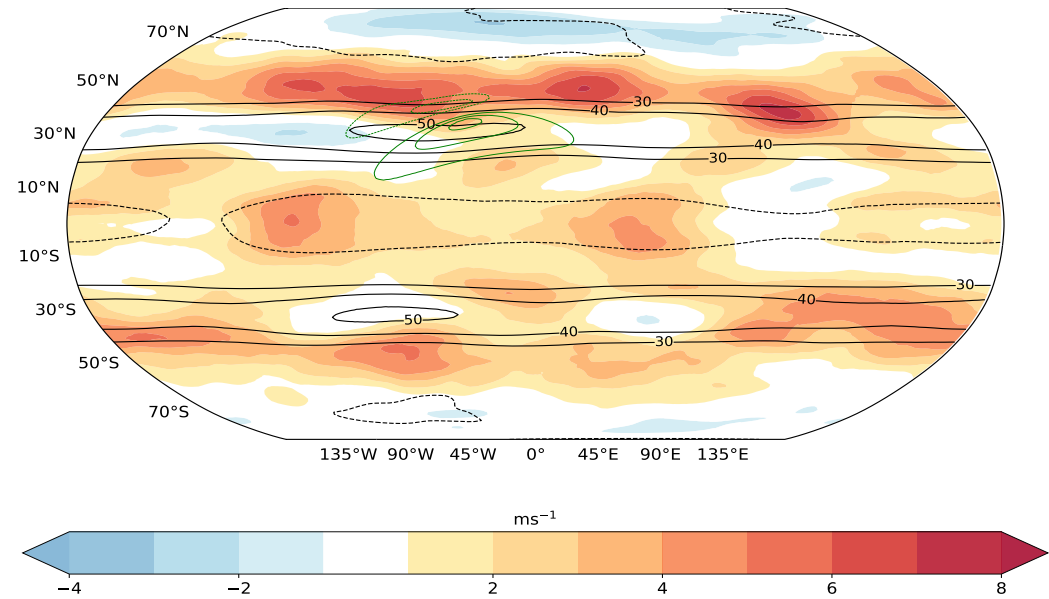
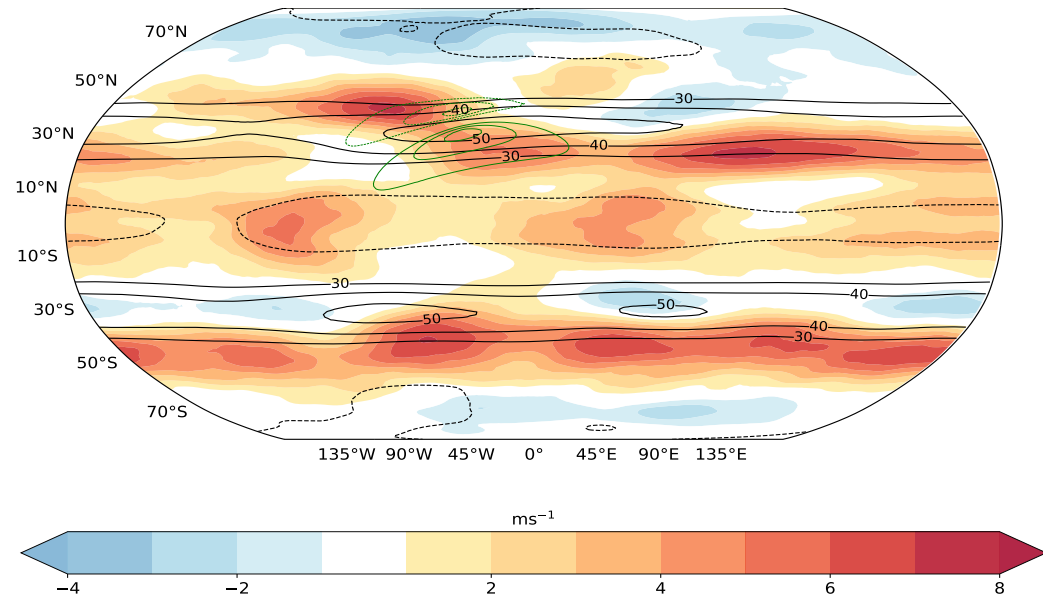


E vector divergence / zonal wind speed difference

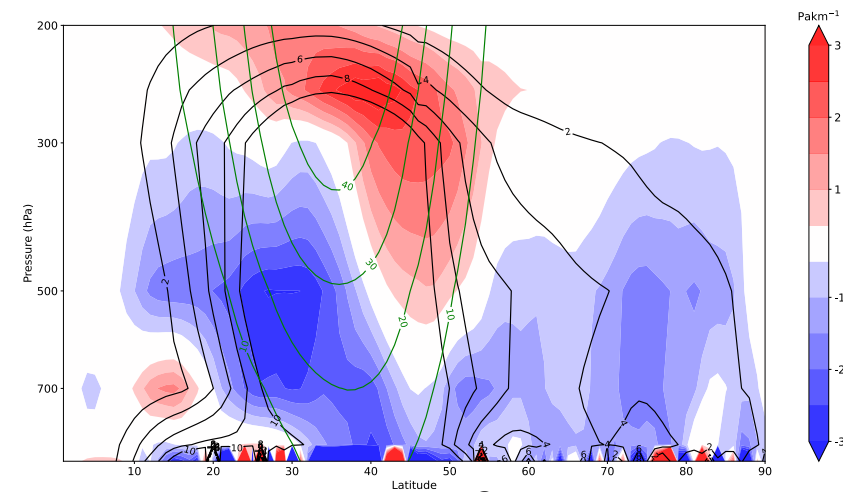


Slope difference/ zonal wind climatology (cntl)

# Sensitivity to SST position



Front at 38N



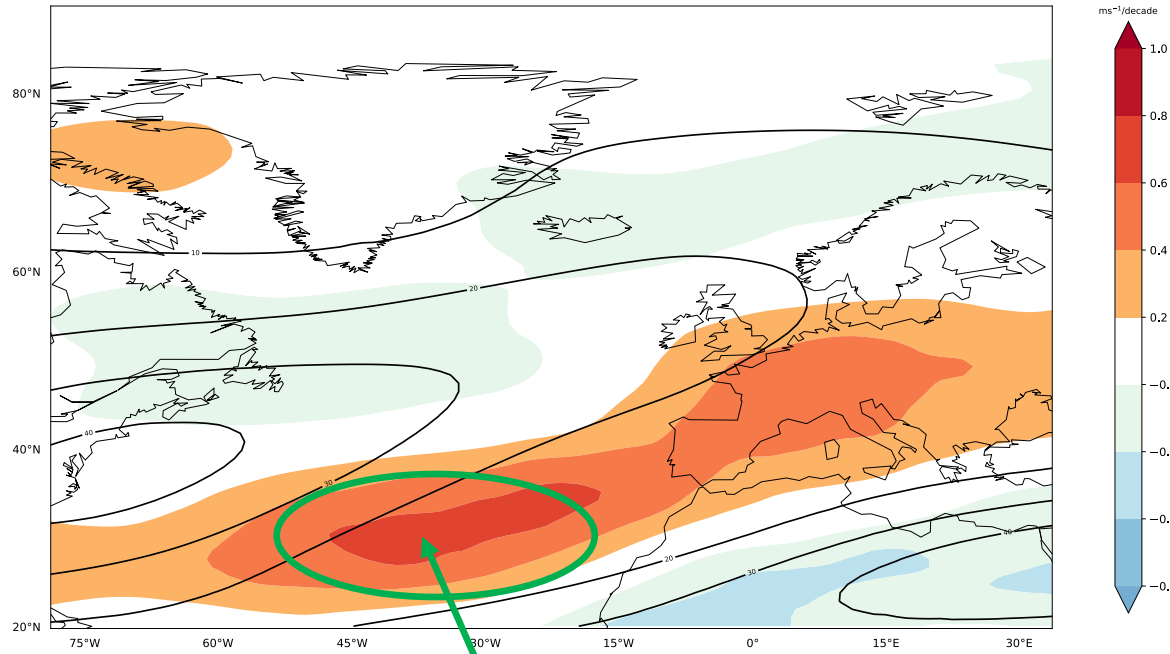
Front at 42N

Fully-coupled CESM simulations

# CESM simulations

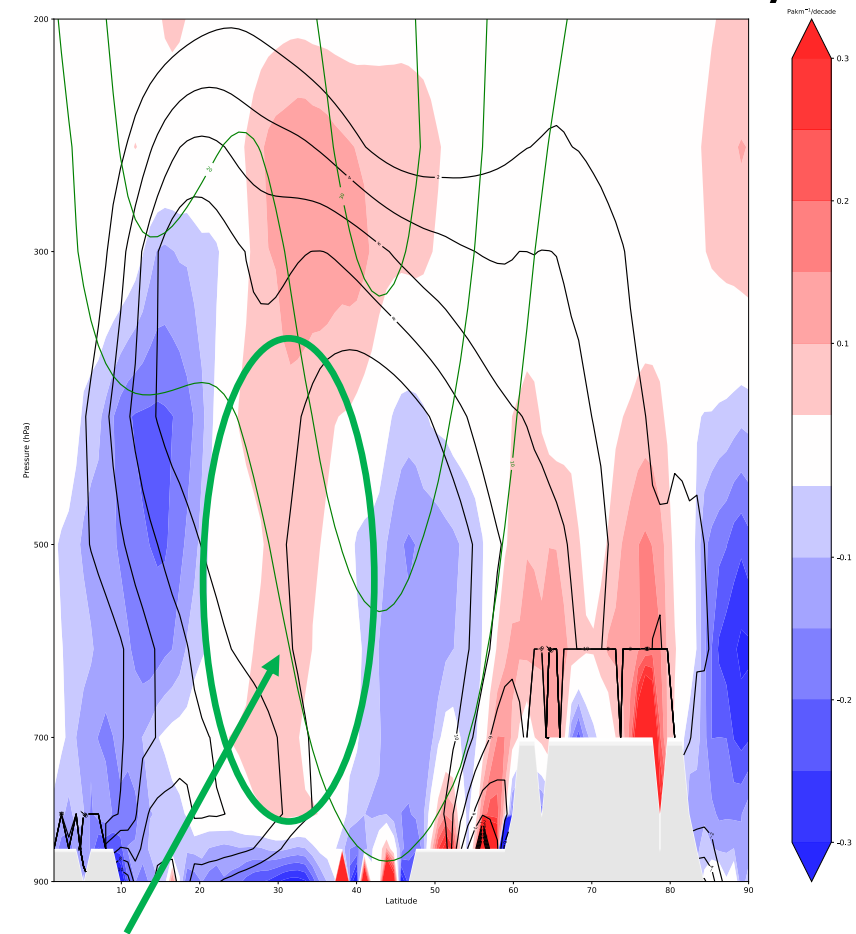
- Community Earth System model version 2.1.2
- Fully-coupled model
- 1° horizontal resolution
- Historical forcing up to 2015 and SSP3-7.0 emissions scenario from 2015 onwards
- 5 ensemble members

## CESM trends (ensemble mean 1980-2022)



Equatorward shift of the jet  
Different to ERA 5 over the Gulf Stream

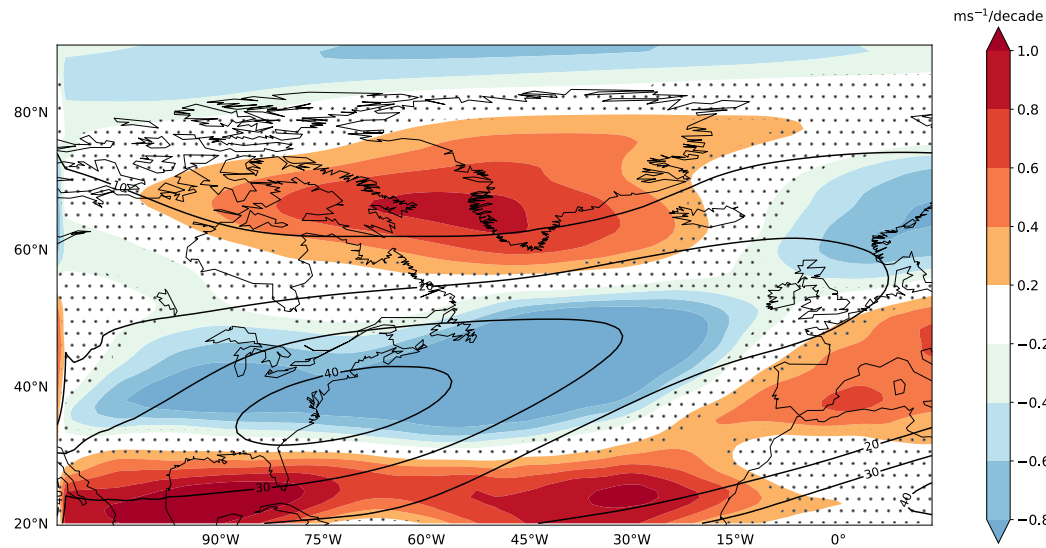
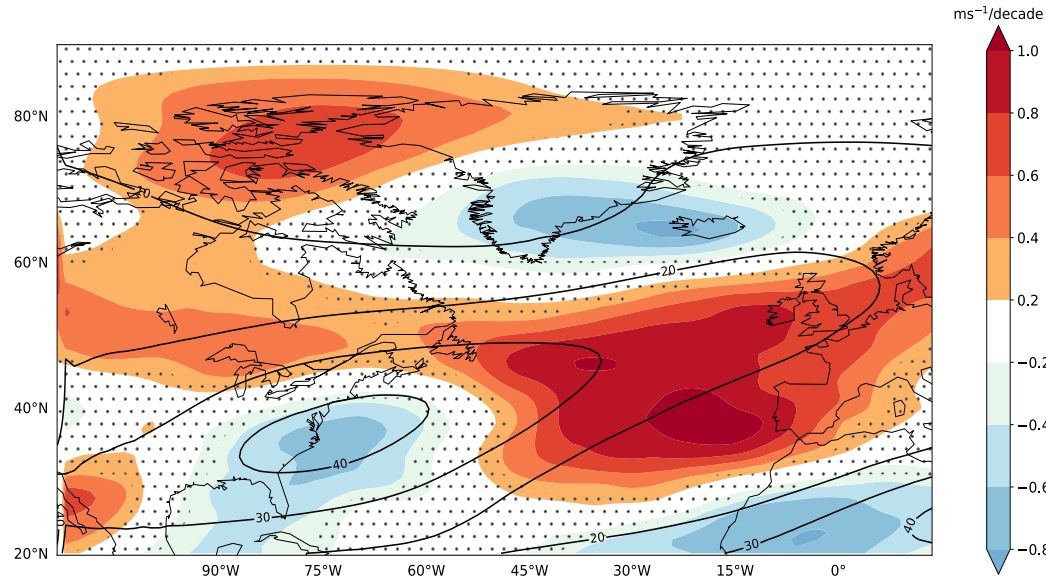
Zonal wind trend at 250 hPa



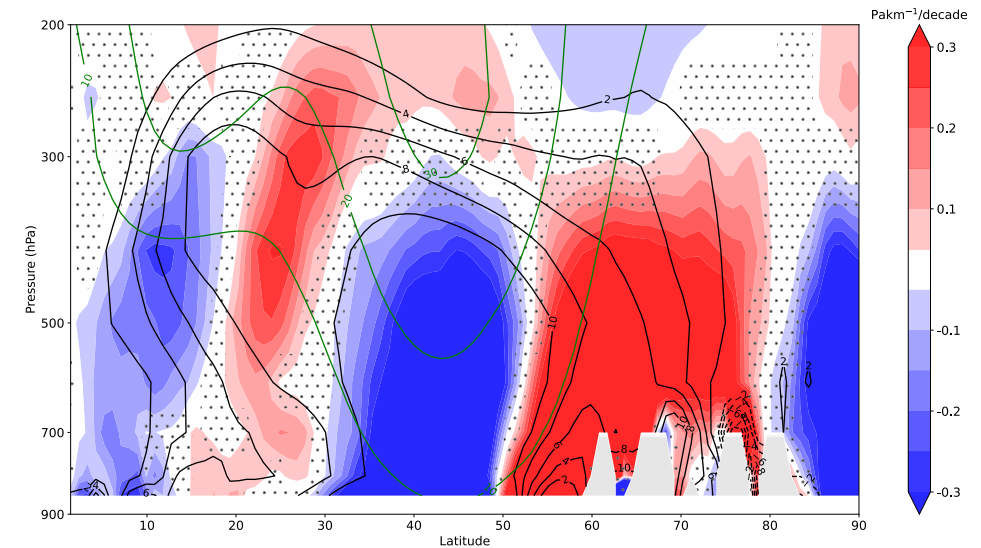
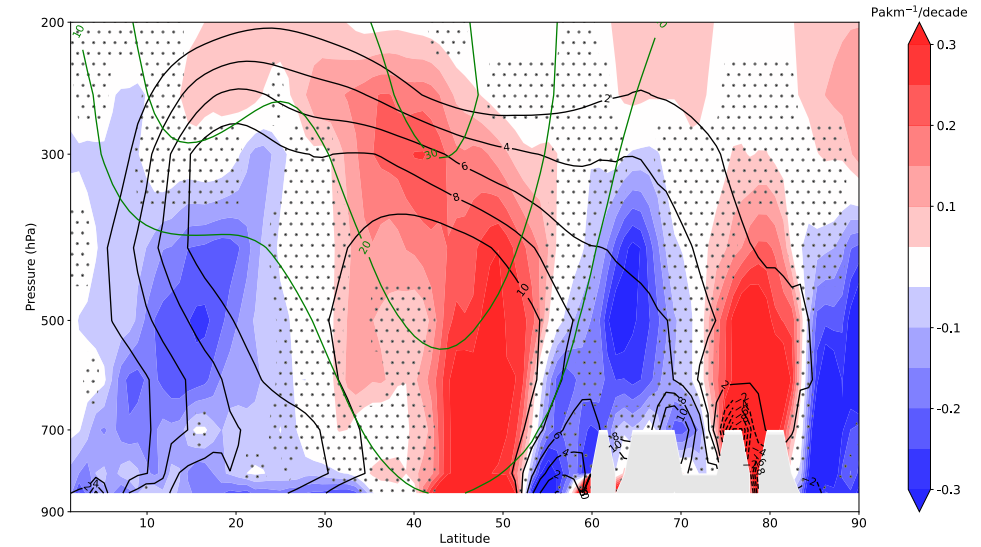
Increase of baroclinicity coincident with positive jet trend

Slope trend (zonal mean 80 – 15 W)

## Some individual members (1980-2022)



Zonal wind trend (250 hPa)



Slope trend (zonal mean 80 – 15 W)



# Conclusions

- The North Atlantic jet stream has slightly shifted poleward over the Gulf Stream and equatorward downstream
- The increase in diabatic seems to be responsible for the local jet trends
- An increase in baroclinicity and eddy momentum convergence accelerates the jet in a zonal band, extending over Europe
- These mechanisms are mostly reproduced in an aquaplanet setup, except the diabatic heating
- Trends in the fully-coupled climate model are different from ERA5 over the Gulf Stream

